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BET4733 Introduction to Coastal Infrastructure

Waves

by

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Chapter Description

- **Expected Outcomes**

Analyze the principles of wave mechanics, tides, littoral processes and coastal sediment transport in methods of shore protection and coastal infrastructures.

- **References**

- 1) Kamphuis, J. William, Introduction to Coastal Engineering and Management, Advanced Series on Ocean Engineering-Volume 30, World Scientific, 2010.
- 2) Reeve D., Chadwick A. and Fleming C. Coastal Engineering-Processes, Theory and Design Practice, CRC Press, 2015.
- 3) Kim Y.C., Design of Coastal Structures and Sea Defences, World Scientific, 2015.
- 4) US Army Corps of Engineers, Coastal Engineering Manual, Washington, 1998-now.



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CONTENTS

- Introduction
- Wave Generation
- Wave Motion
- Fetch Length
- Wave Classification
- Deep Water Wave Movement
- Wave Propagation
- Coastal Engineering Manual, USACE



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INTRODUCTION

WAVES

**ENERGY
SOURCE**

DETERMINE GEOMETRY AND COMPOSITION
OF BEACHES

INFLUENCES PLANNING AND DESIGN OF
COASTAL STRUCTURES



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WAVE GENERATION

- Waves are formed by a disturbing force:
 - ❖ Wind
 - ❖ Underwater earthquake/landslides
 - ❖ Change in atmospheric pressure-seiches
- Restoring forces restore water to its restoring state:
 - ❖ Gravity = gravity waves
 - ❖ Surface Tension = capillary waves



WAVE MOTION

Wave Generation (e.g. wind)

- *seas
- *ripples
- *chops

Wave Propagation

- *swells

Wave Dissipation

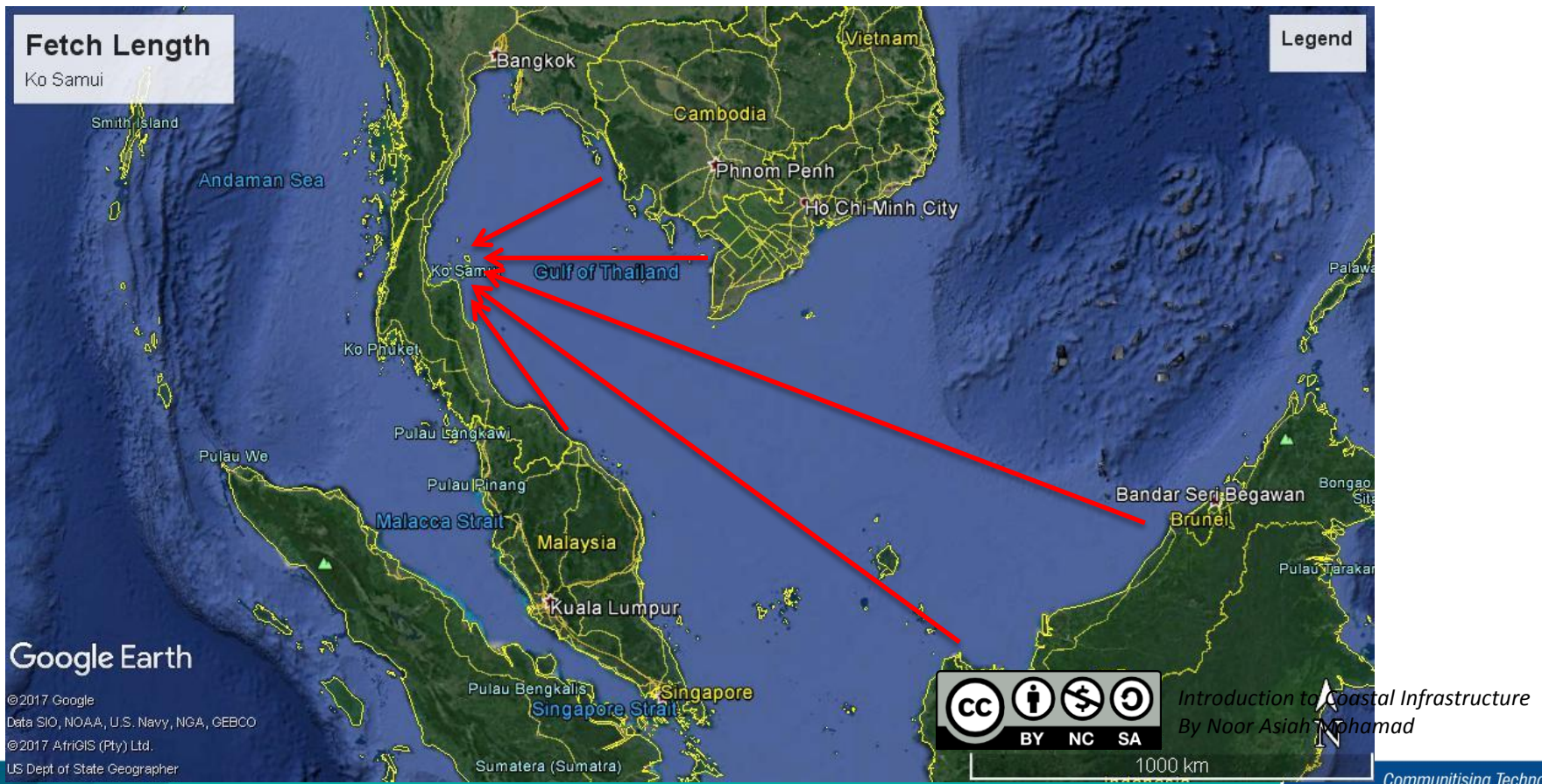
- * wave breakers



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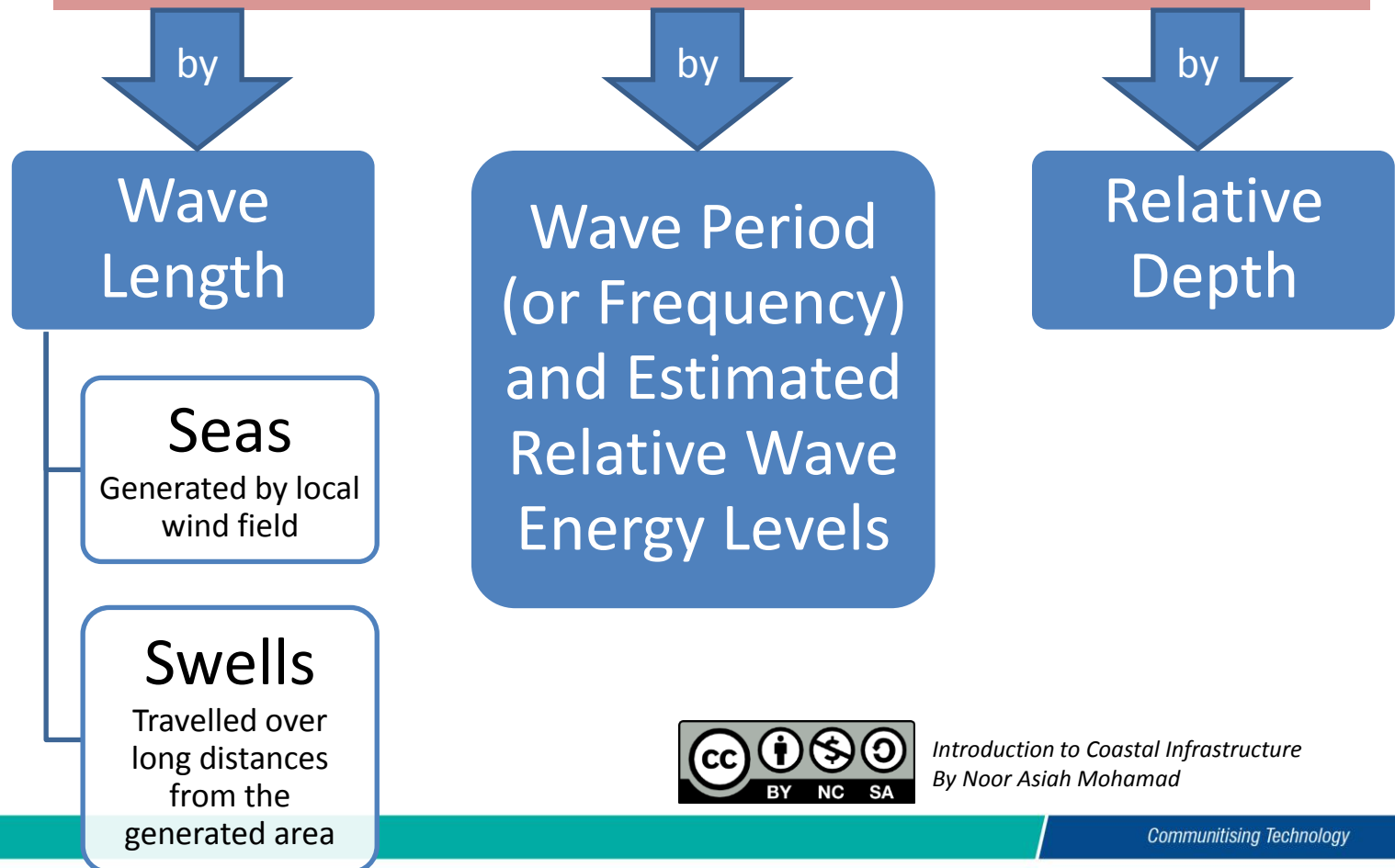
FETCH LENGTH

- Fetch Length: The horizontal distance in the direction of the wind over which a wind generates seas (Glossary of Coastal Terminology, 2012)



WAVE CLASSIFICATION

WAVE CLASSIFICATION



WAVE CLASSIFICATION

Deep-water waves:

$$d/L > 1/2$$

L, v and h remain constant over long distances

Transitional water waves:

$$1/2 > d/L > 1/25$$

L and v decrease, wave height "h" increase rounded tops form peaks

Shallow water waves:

$$d/L < 1/25$$

Particles at top of wave faster than bottom, and wave breaks



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DEEP WATER WAVE MOVEMENT

- In deep water, while wave form moves, water particle below the surface move in stationary circular orbits that get smaller and smaller to a depth of about $0.5 L$.
- In shallow or transitional water, water particle orbits are elliptical.
- H and L change but T remains constant.
- The waves become steeper and steeper and finally breaks in the surf zone.



WAVE PROPAGATION

- Waves can be affected by the following processes as they move from deep to shallow water:
 - ✓ Shoaling – waves becoming higher (energy is compressed over a shorter distance) as they slow down to move from deeper water to shallower water
 - ✓ Refraction - bending effect of wave crest in order to align with bottom contours or bathymetry
 - ✓ Breaking -
 - ✓ Diffraction - wave energy is laterally transferred along a wave crest as the waves bend around an obstruction
 - ✓ Reflection - wave energy is reflected back to sea as the waves hit into a rigid obstruction such as a breakwater, seawall, cliff or a sloping beach



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WAVE BREAKERS

Spilling

Breaking occurs over a distance.

Plunging

Breaking is usually with a crash where smooth splash-up follows

Collapsing

Bubbles & foam are present.

Surging

Little or no bubbles production.



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COASTAL ENGINEERING MANUAL

USACE

<http://www.publications.usace.army.mil/USACE-Publications/Engineer-Manuals/u43544q/636F617374616C20656E67696E656572696E67206D616E75616C/>



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Relative Depth	Shallow Water $\frac{d}{L} < \frac{1}{25}$	Transitional Water $\frac{1}{25} < \frac{d}{L} < \frac{1}{2}$	Deep Water $\frac{d}{L} < \frac{1}{2}$
1. Wave profile	Same As >	$\eta = \frac{H}{2} \cos \left[\frac{2\pi x}{L} - \frac{2\pi t}{T} \right] - \frac{H}{2} \cos \theta$	< Same As
2. Wave celerity	$C = \frac{L}{T} = \sqrt{gd}$	$C = \frac{L}{T} = \frac{gT}{2\pi} \tanh \left(\frac{2\pi d}{L} \right)$	$C = C_0 = \frac{L}{T} = \frac{gT}{2\pi}$
3. Wavelength	$L = T\sqrt{gd} = CT$	$L = \frac{gT^2}{2\pi} \tanh \left(\frac{2\pi d}{L} \right)$	$L = L_0 = \frac{gT^2}{2\pi} = C_0 T$
4. Group velocity	$C_g = C = \sqrt{gd}$	$C_g = nC = \frac{1}{2} \left[1 + \frac{4\pi d/L}{\sinh(4\pi d/L)} \right] C$	$C_g = \frac{1}{2} C = \frac{gT}{4\pi}$
5. Water particle velocity			
(a) Horizontal	$u = \frac{H}{2} \sqrt{\frac{g}{d}} \cos \theta$	$u = \frac{H}{2} \frac{gT}{L} \frac{\cosh[2\pi(z+d)/L]}{\cosh(2\pi d/L)} \cos \theta$	$u = \frac{\pi H}{T} e^{\left(\frac{2\pi z}{L}\right)} \cos \theta$
(b) Vertical	$w = \frac{H\pi}{T} \left(1 + \frac{z}{d} \right) \sin \theta$	$w = \frac{H}{2} \frac{gT}{L} \frac{\sinh[2\pi(z+d)/L]}{\cosh(2\pi d/L)} \sin \theta$	$w = \frac{\pi H}{T} e^{\left(\frac{2\pi z}{L}\right)} \sin \theta$
6. Water particle accelerations			
(a) Horizontal	$a_x = \frac{H\pi}{T} \sqrt{\frac{g}{d}} \sin \theta$	$a_x = \frac{g\pi H}{L} \frac{\cosh[2\pi(z+d)/L]}{\cosh(2\pi d/L)} \sin \theta$	$a_x = 2H \left(\frac{\pi}{T} \right)^2 e^{\left(\frac{2\pi z}{L}\right)} \sin \theta$
(b) Vertical	$a_z = -2H \left(\frac{\pi}{T} \right)^2 \left(1 + \frac{z}{d} \right) \cos \theta$	$a_z = -\frac{g\pi H}{L} \frac{\sinh[2\pi(z+d)/L]}{\cosh(2\pi d/L)} \cos \theta$	$a_z = -2H \left(\frac{\pi}{T} \right)^2 e^{\left(\frac{2\pi z}{L}\right)} \cos \theta$
7. Water particle displacements			
(a) Horizontal	$\xi = -\frac{HT}{4\pi} \sqrt{\frac{g}{d}} \sin \theta$	$\xi = -\frac{H}{2} \frac{\cosh[2\pi(z+d)/L]}{\sinh(2\pi d/L)} \sin \theta$	$\xi = -\frac{H}{2} e^{\left(\frac{2\pi z}{L}\right)} \sin \theta$
(b) Vertical	$\zeta = \frac{H}{2} \left(1 + \frac{z}{d} \right) \cos \theta$	$\zeta = \frac{H}{2} \frac{\sinh[2\pi(z+d)/L]}{\sinh(2\pi d/L)} \cos \theta$	$\zeta = \frac{H}{2} e^{\left(\frac{2\pi z}{L}\right)} \cos \theta$
8. Subsurface pressure	$p = \rho g(\eta - z)$	$p = \rho g \eta \frac{\cosh[2\pi(z+d)/L]}{\cosh(2\pi d/L)} - \rho g z$	



GUIDE FOR USE OF TABLES C1 AND C2

d/L_0 = ratio of depth of water at any specific location to the wavelength in deep water

d/L = ratio of the depth of water at any specific location to the wavelength at that same location

$H/H_0' = K_s$ = ratio of the wave height in shallow water to what its wave height would have been in deep water if unaffected by refraction
= shoaling coefficient = $\sqrt{[1/2(1/n) (1/(C/C_0))]}$

K = a pressure response factor used in connection with underwater pressure instruments, where

$$= \frac{H'}{H} = \frac{P}{P_0} = \frac{\cosh [2\pi d/L (1 + z/d)]}{\cosh (2\pi d/L)} \quad \text{or} \quad \frac{\cosh [2\pi (d + z)/L]}{\cosh (2\pi d/L)}$$

where P is the pressure fluctuation at a depth z measured negatively below still water, P_0 is the surface pressure fluctuation, d is the depth of water from SWL to the ocean bottom, L is the wavelength in any particular depth of water, and H is the corresponding variation of head at a depth z . The values of K shown in the tables are for the instrument placed on the bottom using the equation when

$z = -d = \frac{1}{\cosh (2\pi d/L)}$ values tabulated in Column 8

*Source: Coastal
Engineering Manual,
USACE*



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GUIDE FOR USE OF TABLES C1 AND C2 (cont.)

n = the fraction of wave energy that travels forward with the waveform: i.e. with the wave velocity C rather than the group velocity C_g

$$= \left[\frac{1}{2} \left\{ 1 + \frac{4\pi d/L}{\sinh(4\pi d/L)} \right\} \right] = C_g/C$$

n is also the ratio of group velocity C_g to wave velocity C

C_g/C_o = ratio of group velocity to deepwater wave velocity where

$$C_g/C_o = C_g/C \times C/C_o = n \tanh(2\pi d/L)$$

M = an energy coefficient defined as

$$\frac{\pi^2}{2 \tanh^2(2\pi d/L)}$$

*Source: Coastal
Engineering Manual,
USACE*



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EXAMPLE PROBLEM II-1-1

FIND:

The wave celerities C and lengths L corresponding to depths $d = 200$ meters (656 ft) and $d = 3$ m (9.8 ft).

GIVEN:

A wave with a period $T = 10$ seconds is propagated shoreward over a uniformly sloping shelf from a depth $d = 200$ m (656 ft) to a depth $d = 3$ m (9.8 ft).

SOLUTION:

Using Equation II-1-15,

$$L_0 = \frac{gT^2}{2\pi} = \frac{9.8 T^2}{2\pi} = 1.56 T^2 \text{ m (5.12 } T^2 \text{ ft)}$$

$$L_0 = 1.56T^2 = 1.56(10)^2 = 156 \text{ m (512 ft)}$$

For $d = 200$ m

$$\frac{d}{L_0} = \frac{200}{156} = 1.2821$$

Note that for values of

$$\frac{d}{L_0} > 1.0$$

$$\frac{d}{L_0} = \frac{d}{L}$$

therefore,

$$L = L_0 = 156 \text{ m (512 ft)} \text{ (deepwater wave, since } \frac{d}{L} > \frac{1}{2} \text{)}$$

which is in agreement with Figure II-1-5.

By Equation II-1-7

$$C = \frac{L}{T} = \frac{156}{T}$$

$$C = \frac{156}{10} = 15.6 \text{ m/s (51.2 ft/s)}$$

For $d = 3$ m

$$\frac{d}{L_0} = \frac{3}{156} = 0.0192$$

Source: Coastal Engineering Manual, USACE

Example Problem II-1-1 (Continued)



Example Problem II-1-1 (Concluded)

By trial-and-error solution (Equation II-1-21) with d/L_0 , it is found that

$$\frac{d}{L} = 0.05641$$

hence

$$L = \frac{3}{0.05641} = 53.2 \text{ m (174 ft)} \left(\text{transitional depth, since } \frac{1}{25} < \frac{d}{L} < \frac{1}{2} \right)$$

$$C = \frac{L}{T} = \frac{53.2}{10} = 5.32 \text{ m/s (17.4 ft/s)}$$

An approximate value of L can also be found by using Equation II-1-11

$$L = \frac{gT^2}{2\pi} \sqrt{\tanh\left(\frac{4\pi^2 d}{T^2 g}\right)}$$

which can be written in terms of L_0 as

$$L = L_0 \sqrt{\tanh\left(\frac{2\pi d}{L_0}\right)}$$

therefore

$$L = 156 \sqrt{\tanh\left(\frac{2\pi(3)}{156}\right)}$$

$$L = 156 \sqrt{\tanh(0.1208)}$$

$$L = 156 \sqrt{0.1202} = 54.1 \text{ m (177.5 ft)}$$

which compares with $L = 53.3$ m obtained using Equations II-1-8, II-1-9, or II-1-21. The error in this case is 1.5 percent. Note that Figure II-1-5 or Plate C-1 (SPM 1984) could also have been used to determine d/L .

Source: Coastal
Engineering
Manual, USACE



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